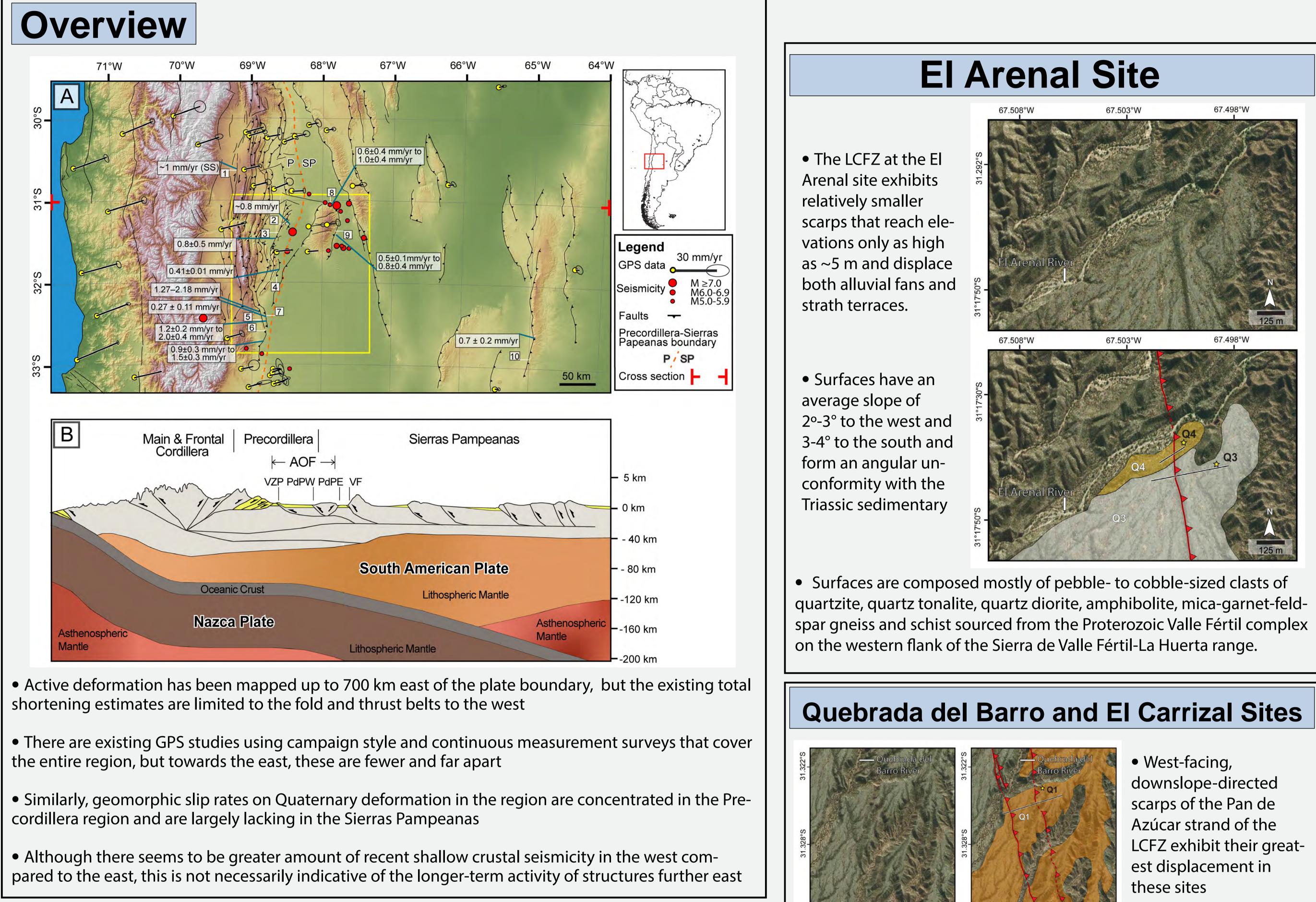
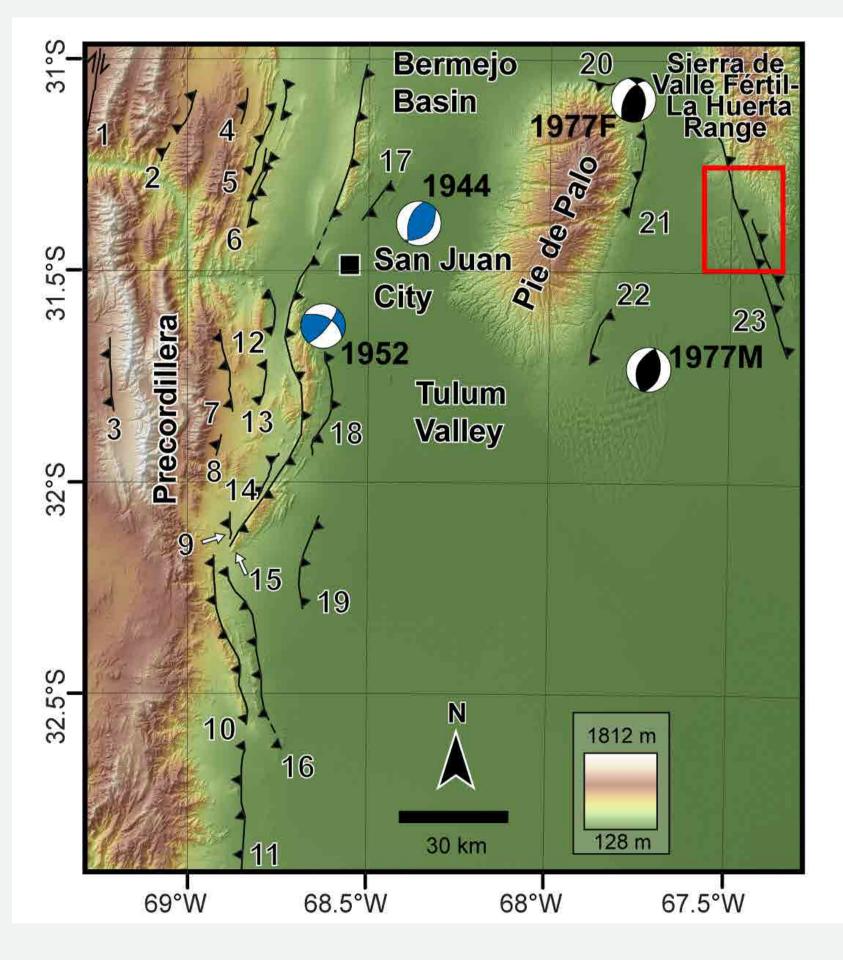


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Background



• Quaternary Andean backarc deformation within the Pampean flat-slab segment (27–33.5°S) is widely distributed, extending up to 700 km east from the Peru-Chile trench.

• In the past 10–15 Ma, up to ~75% of the total shortening between the plate boundary and the forearc region at 30°S has been accommodated in the thin-skinned orogenic belts of the Precordillera.

• Activity in the easternmost Precordillera that commenced at ~2.6 Ma and in the westernmost Sierras Pampeanas at ~3 Ma focused deformation between these two sections of the retroarc between 30 and 32°S, forming the Andean Orogenic Front.

• Compared to the thin-skinned structures in the west, little is known about the rates of movement of most of the basement-cored uplifts of the Sierras Pampeanas.

• GPS studies demonstrate an overall west-east gradient of decreasing shortening rates

• It is unclear, however, if this trend holds true on the longer term and if it resembles the trend in level of activity of structures, due to the sparseness of observation points, poor vertical positional accuracy, and limited temporal coverage of GPS studies.

Late Quaternary Intraplate Deformation: The Las Chacras Fault Zone, West-Central Argentina

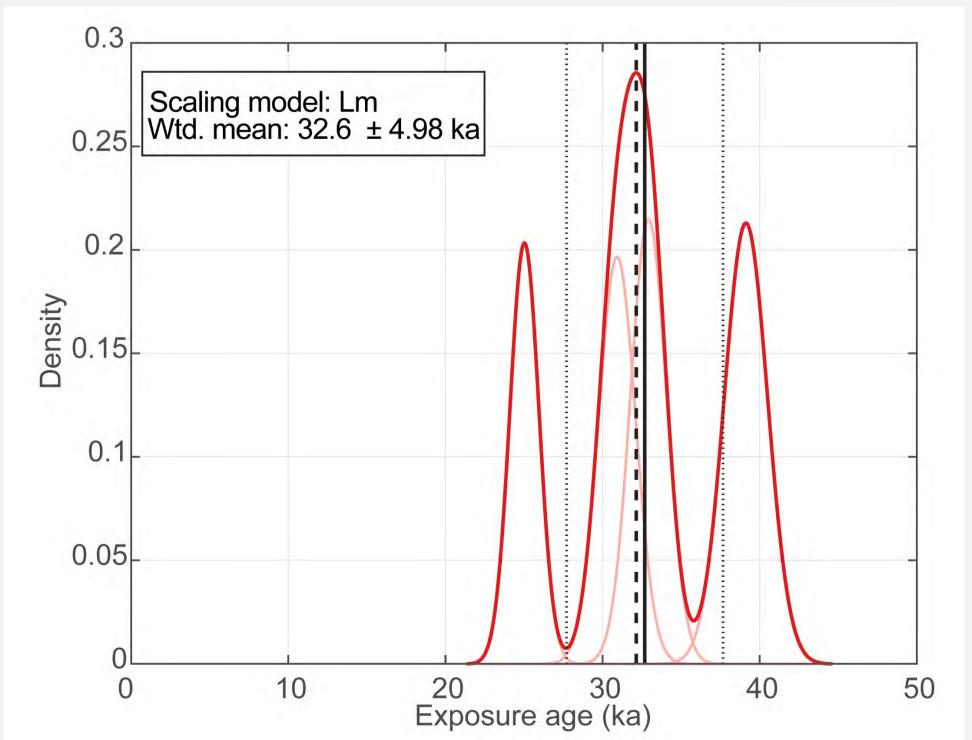
Jeremy Rimando^{1,2}, Lindsay Schoenbohm^{1,2}, Gustavo Ortiz³, Patricia Alvarado³, Agostina Venerdini³, Lewis Owen⁴ Erin Seagren^{1,2}, Paula Marques Figueiredo⁴, Sarah Hammer⁵

• ~20-m-high scarps dis place well preserved late Quaternary alluvial

 Trends N15°W and bifurcates towards the north, with the western trace being more continuous than the eastern trace.

Cosmogenic dating

• **Q4**: 10.0 ± 1.2 ka.



• Q2: 38.7 ± 3.7 ka.; Q2SA, Q2SB, Q2SC and Q2SD yield similar ages of 32.9 \pm 2.6, 25.0 \pm 2.0, 39.1 \pm 2.1, and 30.9 \pm 2.5 ka, respectively (weighted aver age age of 32.5 ± 5.0 ka). • **Q3**: 19.7 ± 4.8 ka.

Quebrada del Barro & El Carrizal Legend Holocene Sediments Quaternary alluvial deposits (Q3) Quaternary alluvial deposits (Q2) Quaternary alluvial deposits (Q1) Triassic Quebrada del Barro Fm. Triassic Carrizal Fm. Triassic Esquina Colorada Fm. Proterozoic Valle Fertil Complex Villarcan Gneiss **Quartz Tonalite** Quartz Diorite 67.58°W 67.50°W

 The LCFZ, is a NW- trending, ~30-km-long, E-dipping zone of reverse faults that are aligned to the southern end of the ~180-km-long Valle Fértil Fault, which bounds the the Sierra de Valle Fértil-La Huerta Range (composed of a Proterozoic igneous and metamorphic basement complex overlain by Triassic sedimentary rocks).

 It is located ~100 km east of the city of San Juan in Argentina. Two main fault strands comprise the LCFZ, the Pan de Azúcar Fault in the west and the Rickard Fault in the east.

 Ages were computed from amalgamated clast samples. Inheritance was accounted for by subtracting the 10Be concentration of adjacent modern channel clasts

• A kernel density estimate (KDE) of individual clast ages (4 clasts: Q2SA, Q2SB, Q2SC, Q2SD) demonstrate good internal consistency.

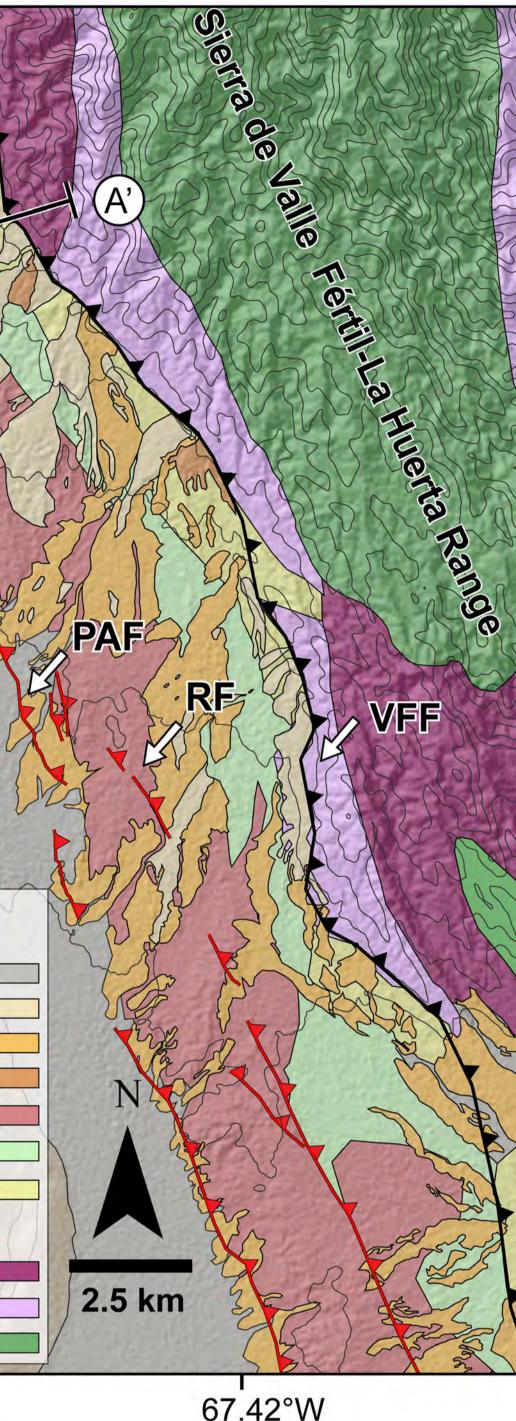
• Due to the ubiquity of desert varnish on all surfaces, which is an indication of high degree of surface preservation, clast ages were computed assuming zero erosion.

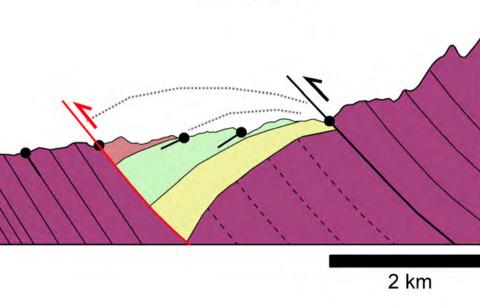
• A sensitivity test was conducted to test effect of erosion rates on age.

• 1 mm/ka of erosion for Q4, Q3, and Q2 underestimates ages by 1, 2 and 3% of the zero-erosion ages, respectively. 5 mm/ka erosion rate underestimates ages by 4, 8, and 20% of the zero-erosion ages, respectively.

 For large clasts, cosmogenic ages are likely to be overestimated due to the inability to fully account for inheritance rather than underestimated due to erosion.

Las Chacras Fault Zone





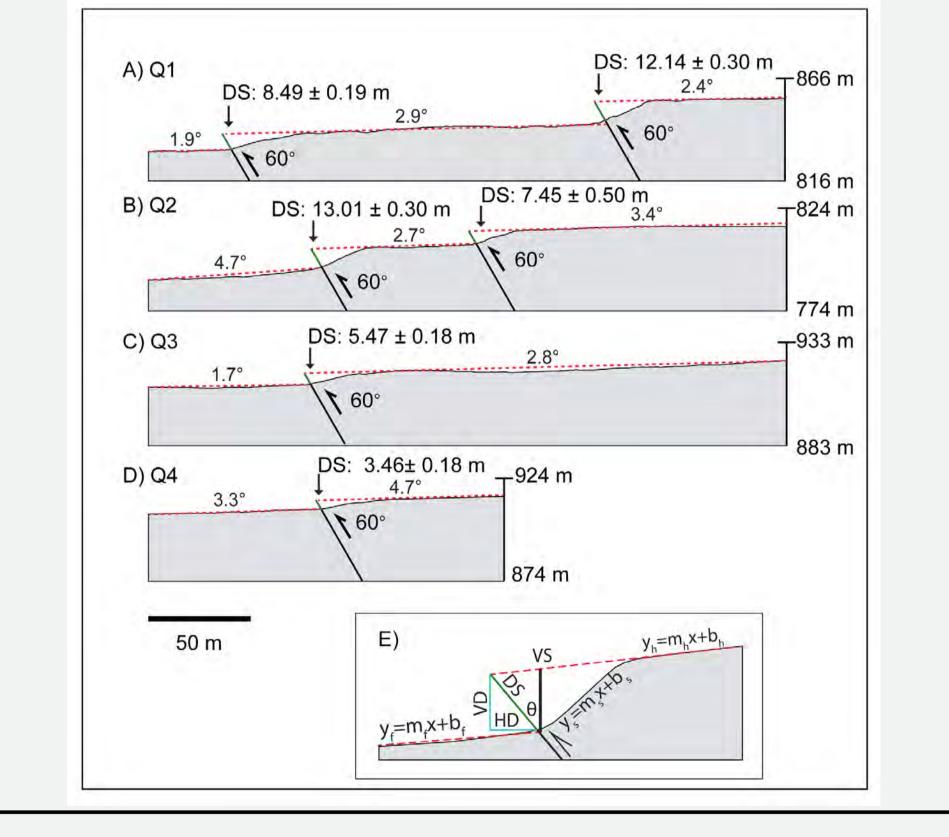
Cumulative Scarps

Profile views of the Pan de Azúcar strand of the LCFZ. A. Scarps at the El Carrizal site where Q2 is displaced. B. Scarps at the Quebrada del Barro site where Q1 is displaced.

Calculation of Displacement

• DGPS scarp profiles were used as markers to compute displacement (dip-slip, shortening, and uplift) according to the methodology described in Yang et al. (2015).

• Given the orientation of the LCFZ, a left-lateral component is likely. However, it is difficult to quantify this due to the absence of preserved, well-defined piercing points that will allow distinguishing the amounts of separation (a.k.a. apparent slip) from actual lateral-slip on ephemeral streams that cut across the fault scarp diagonally.



Conclusions

• We determined geomorphically-derived average shortening and uplift rates for the LCFZ of 0.19 ± 0.02 and 0.33 ± 0.03 mm/yr, respectively, from 10.0 ± 1.2 ka to 38.7 ± 3.7 ka.

• Although faults in the Sierras Pampeanas region still widely lack slip rates, the shortening rate we measure on the LCFZ is significantly lower than on most faults to its west both in the Precordillera and in the Sierra Pie de Palo of the western Sierras Pampeanas; this clearly supports a general west-east decreasing trend of shortening rates across the Pampean flat slab.

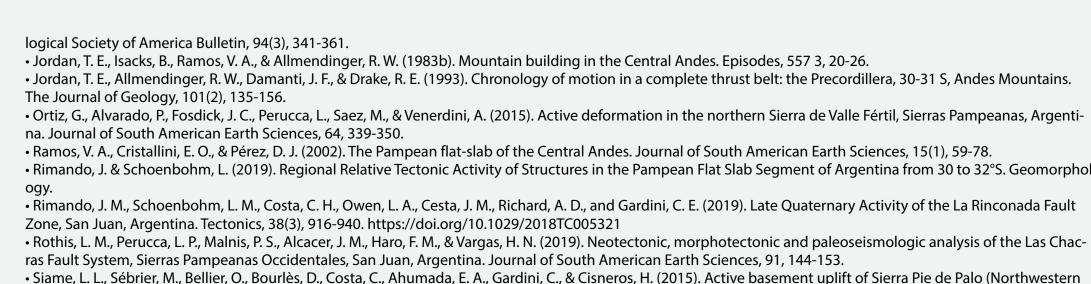
 Further detailed analysis of event horizons from paleoseismic trenching are needed to provide further constraints on timing and recurrence of earthquakes associated with this fault. • The LCFZ, with its abruptly lower shortening rate, also coincides spatially with the resumption to a more steeply-dipping Nazca plate between 67°W and 68°W. Uplift rates, on the other

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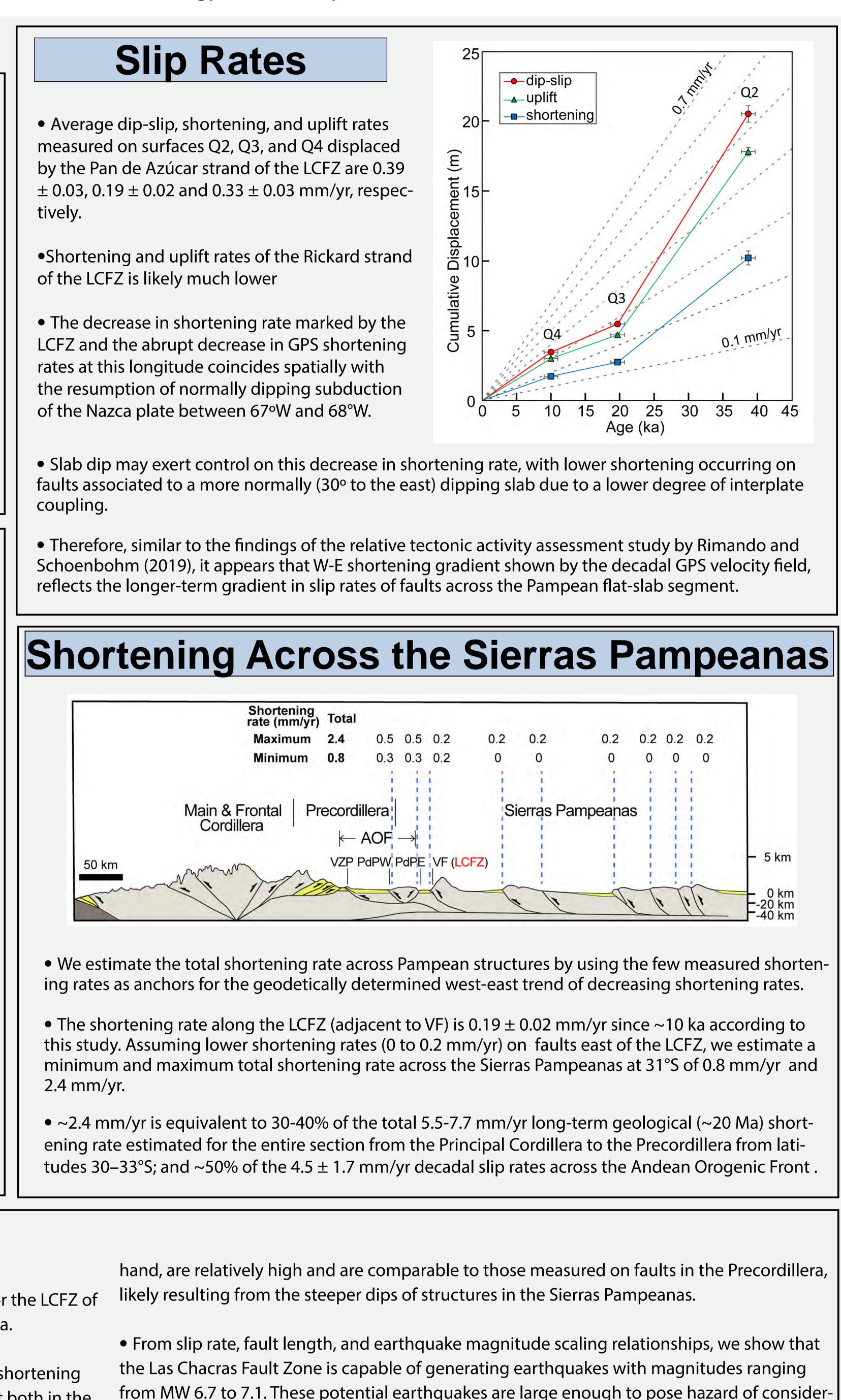
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from MW 6.7 to 7.1. These potential earthquakes are large enough to pose hazard of considerable ground-shaking and liquefaction-related damage to nearby cities.

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